

# Investigating large-scale atmospheric phenomena using complex networks and nonlinear time series analysis tools

Cristina Masoller

Departamento de Física

Universitat Politècnica de Catalunya



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BARCELONATECH

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[cristina.masoller@upc.edu](mailto:cristina.masoller@upc.edu)



[@cristinamasoll1](https://twitter.com/cristinamasoll1)

# THANKS TO

Ignacio Deza    Giulio Tirabassi    Dario Zappala    Riccardo Silini    Marcelo Barreiro



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[cristina.masoller@upc.edu](mailto:cristina.masoller@upc.edu)



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# **Methods for climatological time series analysis**

# Outline

- Univariate analysis: Hilbert analysis

$$\{x_1, x_2, \dots, x_N\}$$

- Bivariate analysis: Correlation and causality measures

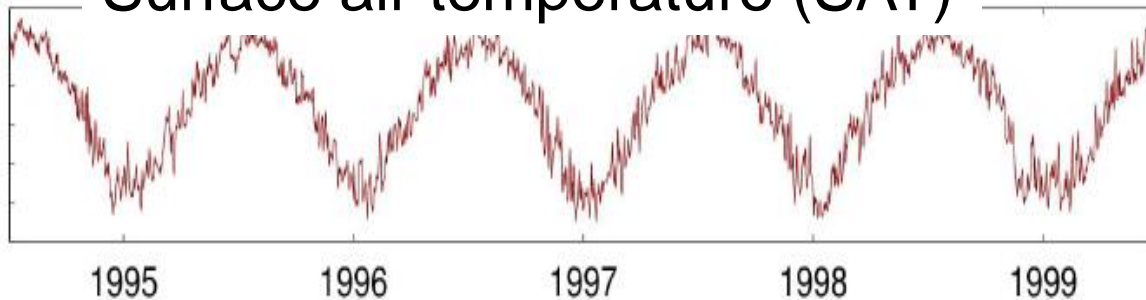
$$\{x_1, x_2, \dots, x_N\} \quad \{y_1, y_2, \dots, y_N\}$$

- Multivariate analysis: Ordinal analysis and complex networks

# The Hilbert Transform (HT)

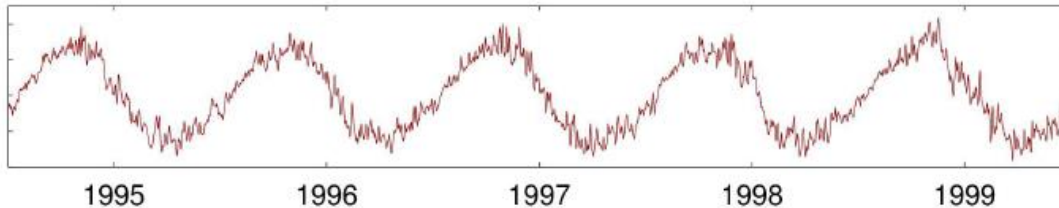
Surface air temperature (SAT)

x

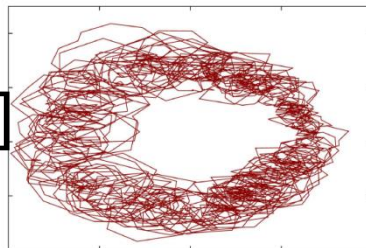


$$\text{HT}[\sin(\omega t)] = \cos(\omega t)$$

HT[x]



y=HT[x]



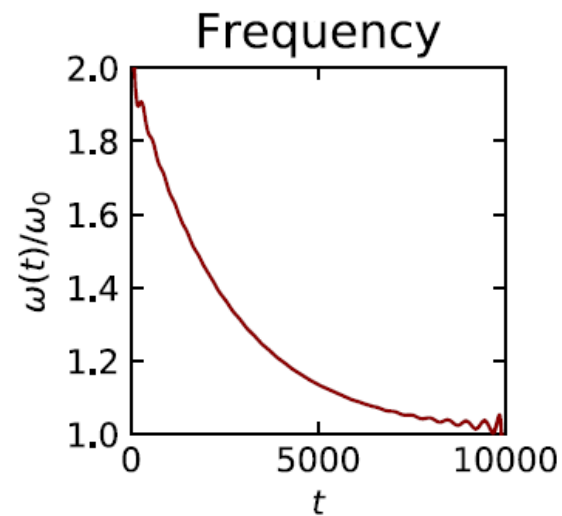
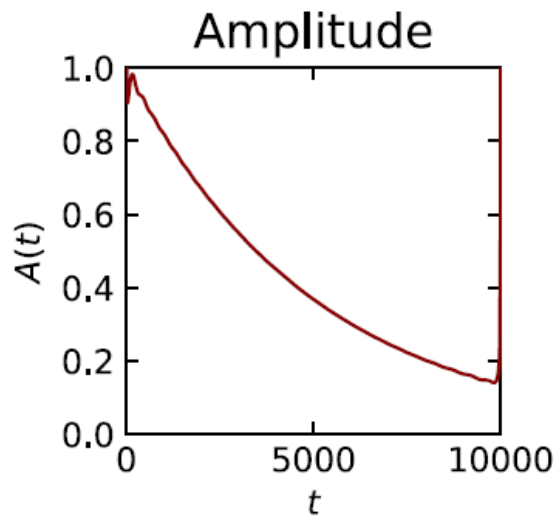
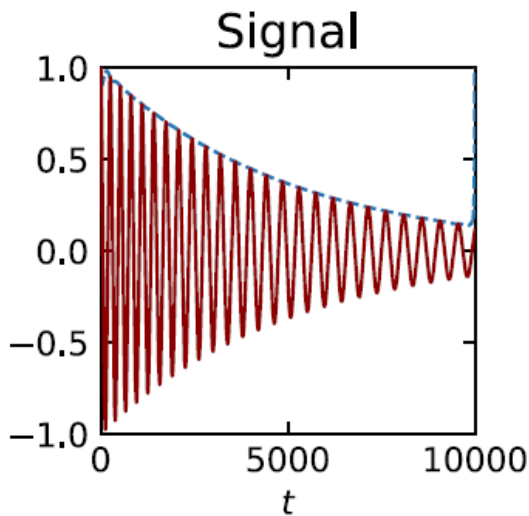
x

$$a(t) = \sqrt{[x(t)]^2 + [y(t)]^2}$$
$$\varphi(t) = \arctan[y(t)/x(t)]$$

$$\omega(t) = d\varphi/dt$$

# Example

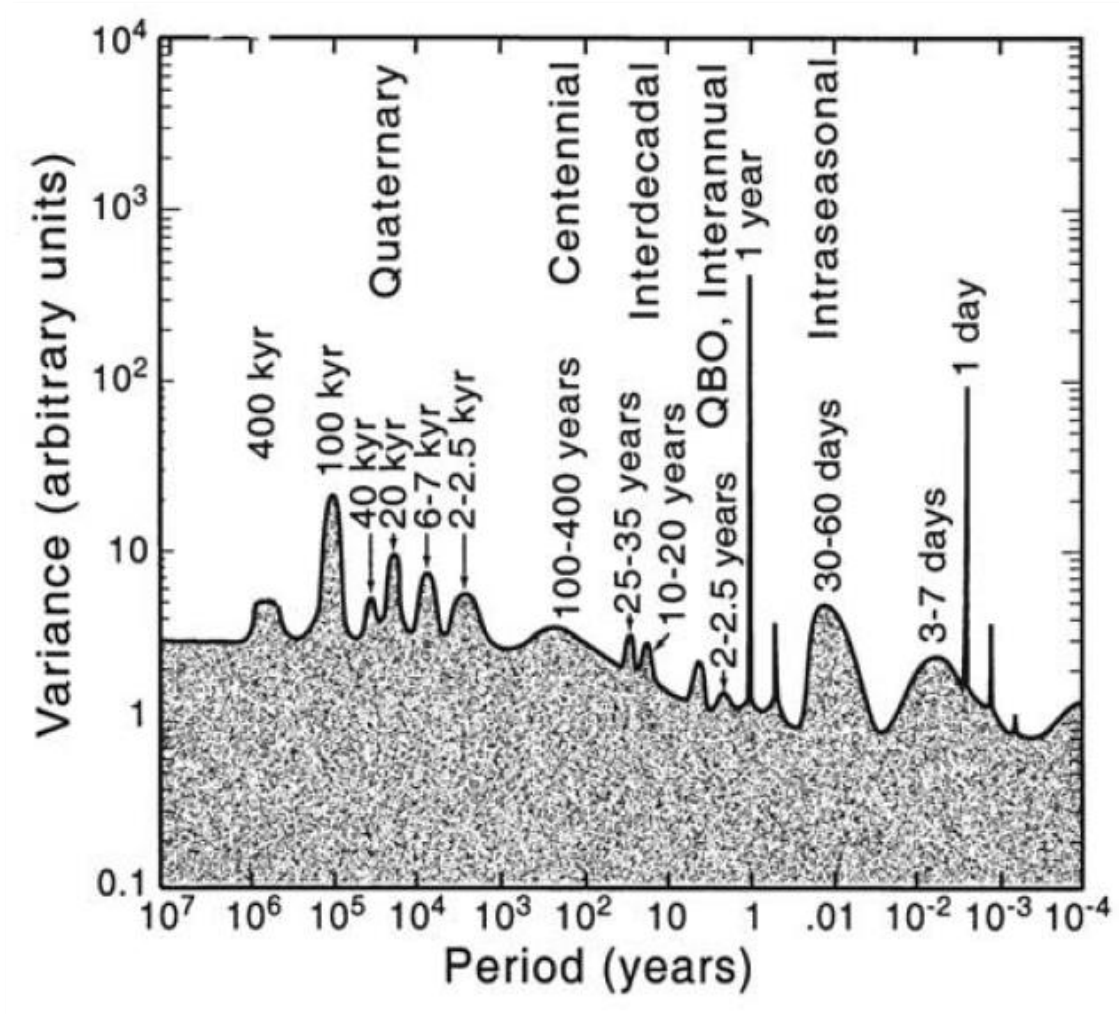
$$x(t) = e^{-\alpha t} \cos \left[ \left( 1 + e^{-2\alpha t} \right) \omega_0 t \right].$$



A word of warning: only if  $x(t)$  is a “narrow-band” signal then  $a(t)$  and  $\omega(t) = d\phi/dt$  have clear physical meaning

- $a(t)$  is the envelope of  $x(t)$
- $\omega(t)$  is the main frequency in the Fourier spectrum

# PROBLEM ! Climatic time series are NOT narrow-band.



An “artist’s representation” of the power spectrum of climate variability (M. Ghil 2002).

## Solution ?

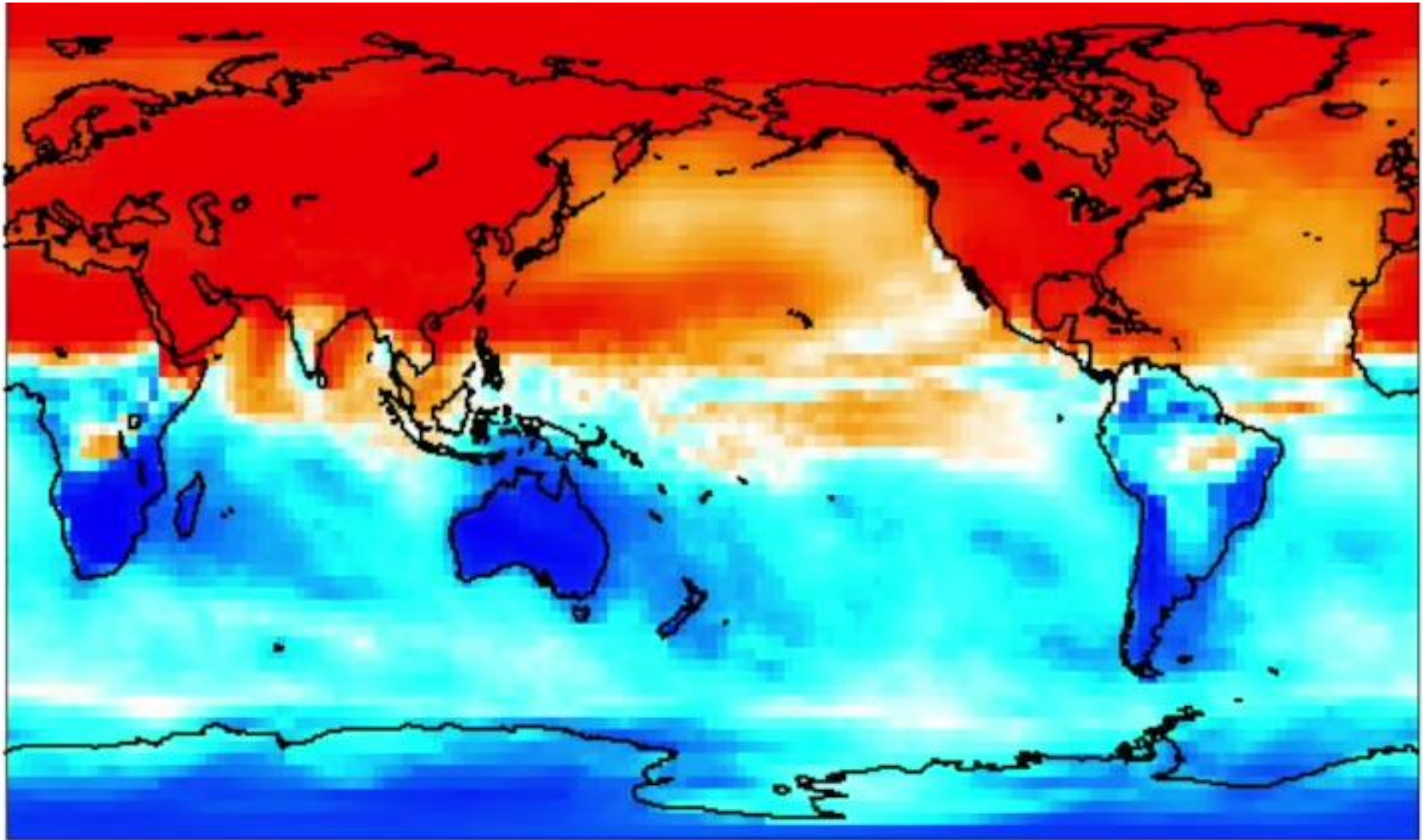
- Isolate a narrow frequency band (usual for EEG analysis).
- However, I will show that HT directly applied to raw surface air temperature (SAT) returns meaningful results.

## Surface Air Temperature (SAT) data

- Spatial grid  $2.5^{\circ} \times 2.5^{\circ} \Rightarrow 10226$  time series
- Daily resolution **1979 – 2016**  $\Rightarrow$  each time series has 13700 data points

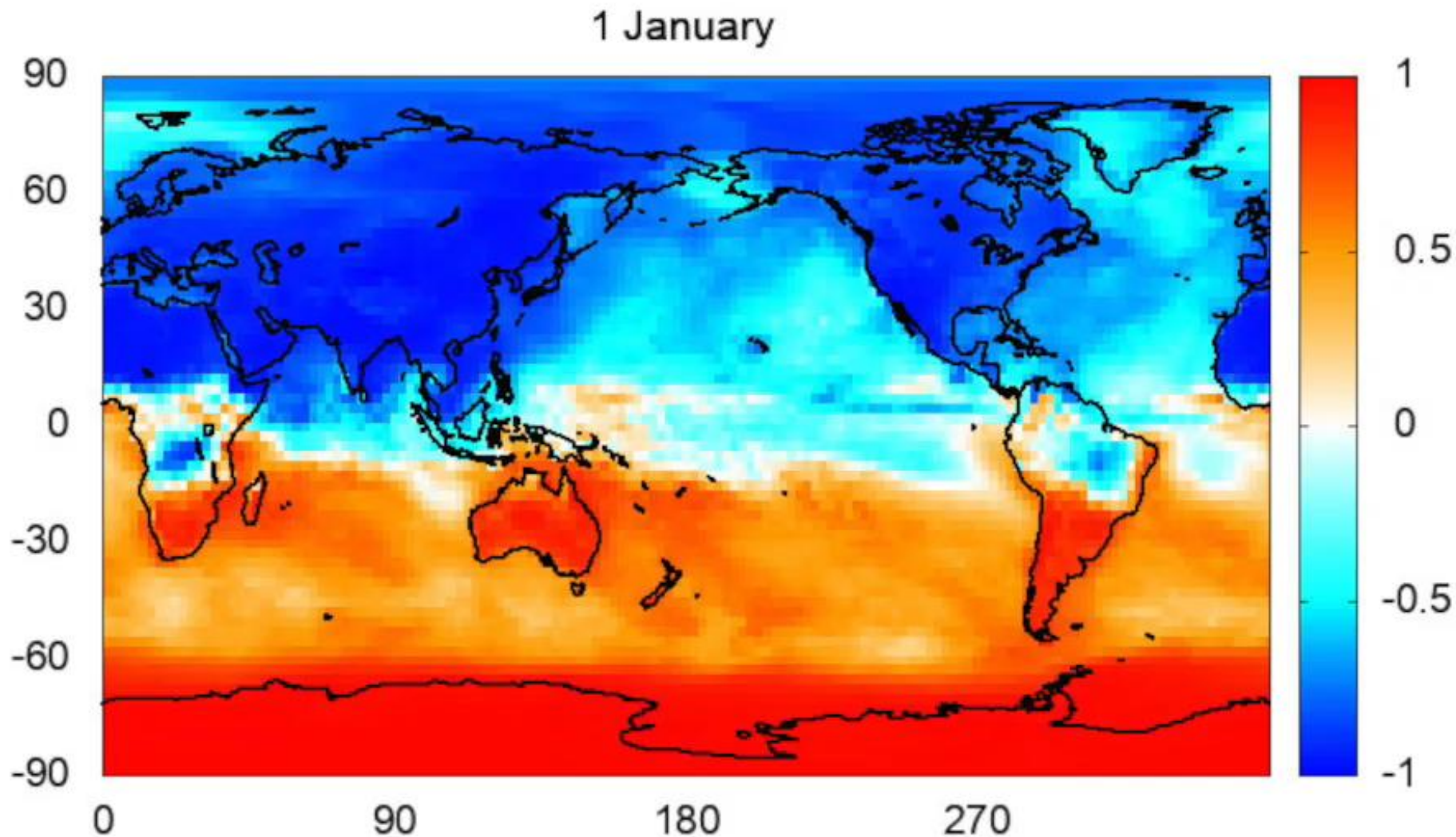
# Which information carries the Hilbert phase?

In color code the  $\cos(\varphi)$  averaged over all **July 1** in the period 1979 – 2016.

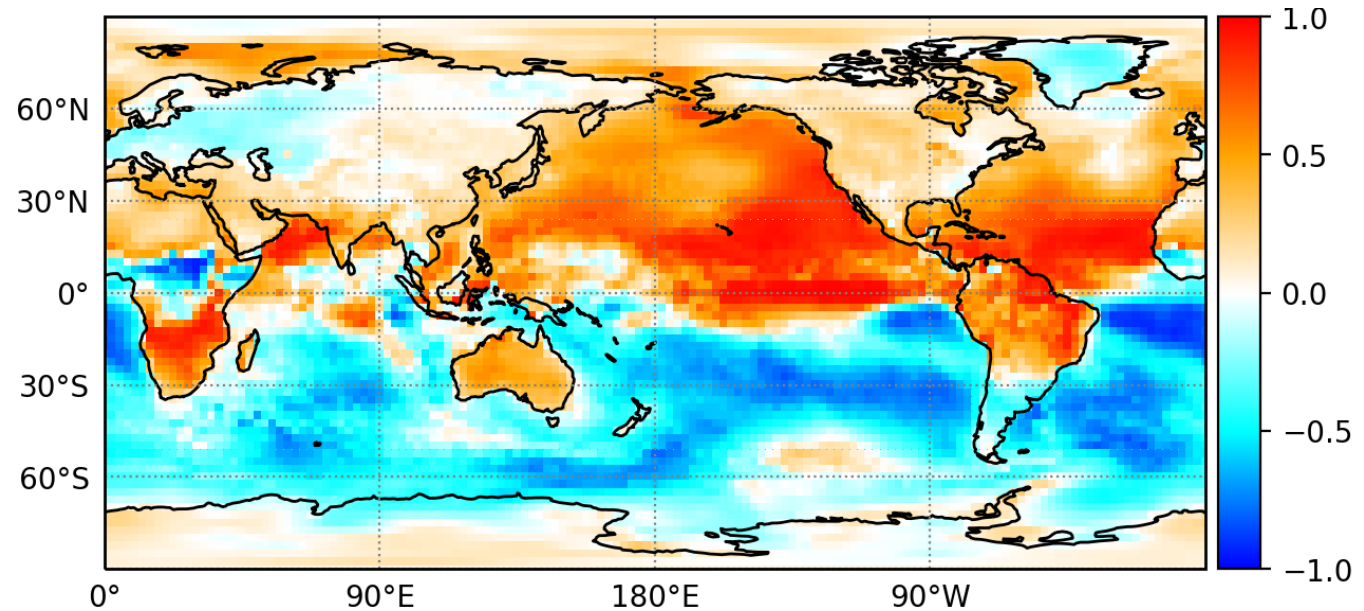


# How do the seasons evolve?

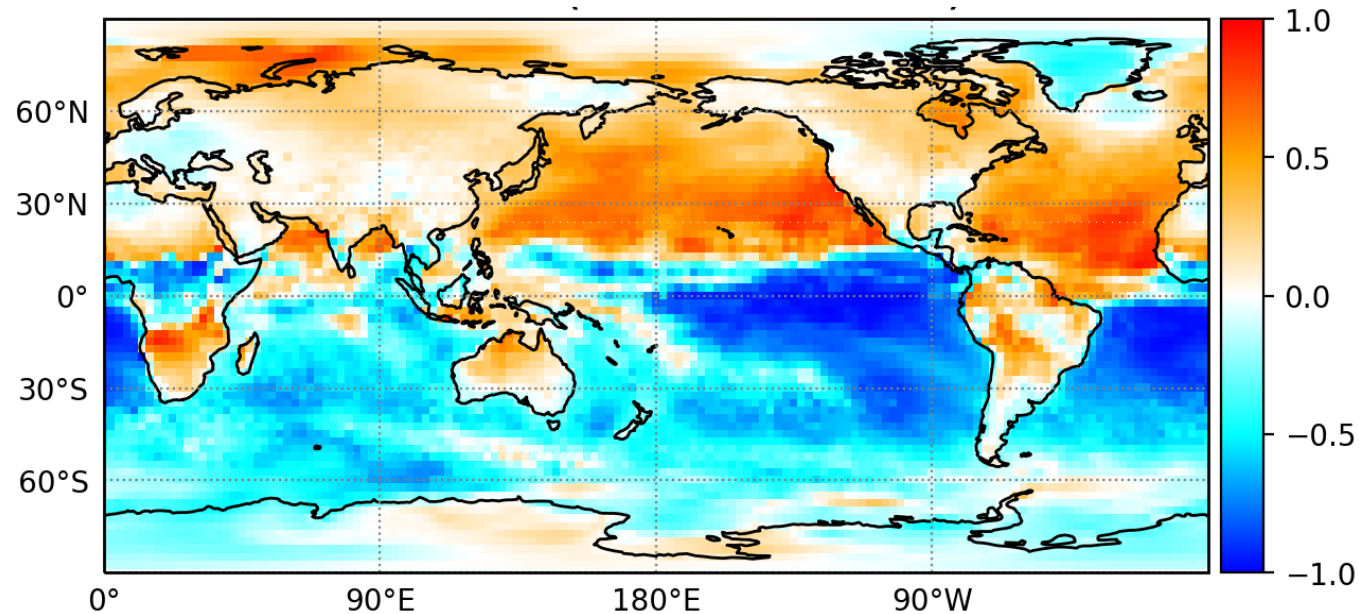
## Temporal evolution of the cosine of the Hilbert phase



El Niño period  
(October 2015)



La Niña period  
(Octubre 2011)



**Can we use Hilbert analysis  
to identify and quantify  
regional “climate change”?**

## Relative decadal variations in each region (“node”)

$$\Delta a = \langle a \rangle_{2016-2007} - \langle a \rangle_{1988-1979}$$

$$\frac{\Delta a}{\langle a \rangle_{2016-1979}}$$

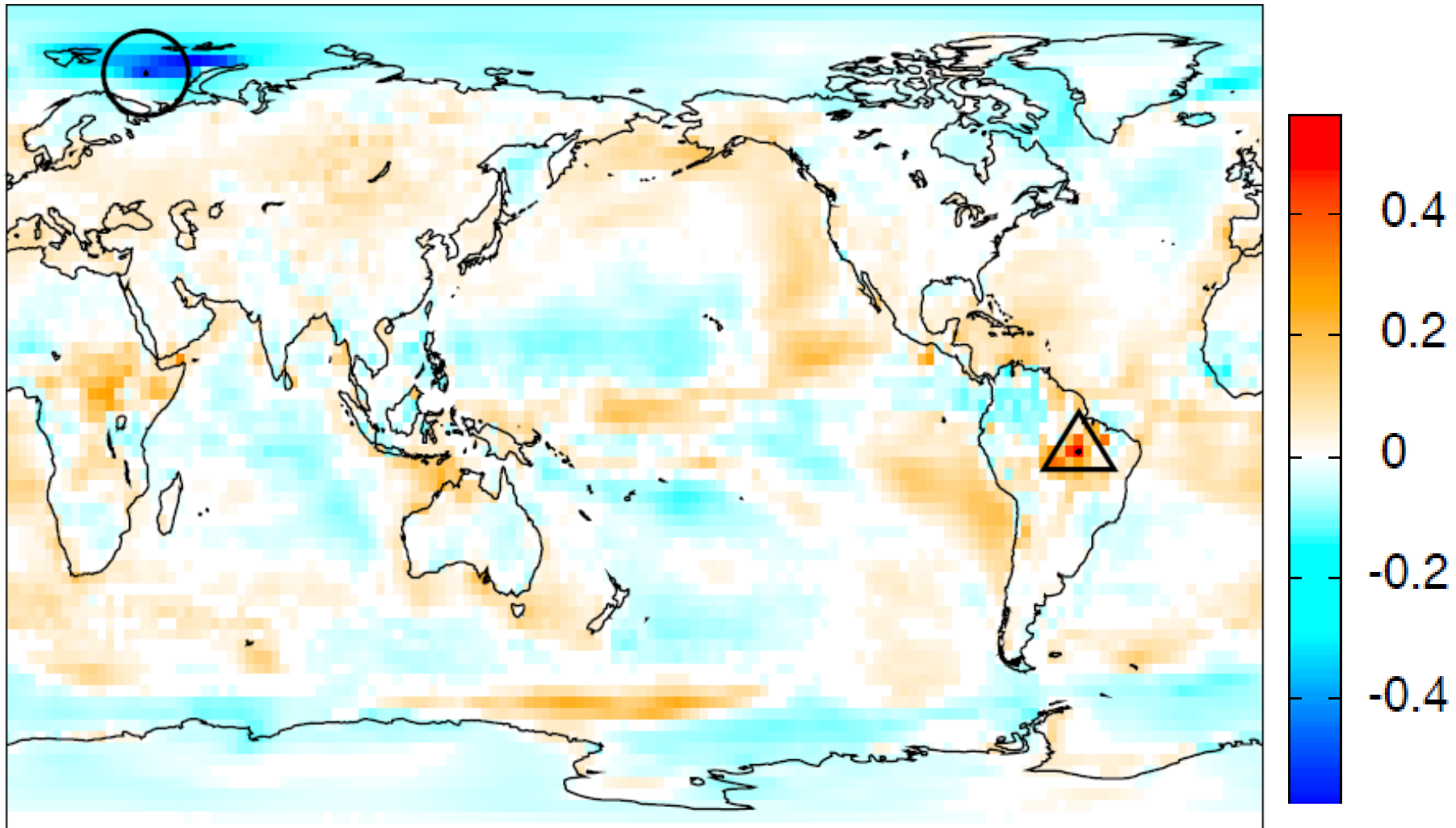
Relative variation is considered **significant** if:

$$\frac{\Delta a}{\langle a \rangle} \geq \langle . \rangle_s + 2\sigma_s \quad \text{or} \quad \frac{\Delta a}{\langle a \rangle} \leq \langle . \rangle_s - 2\sigma_s$$

100 “surrogates”

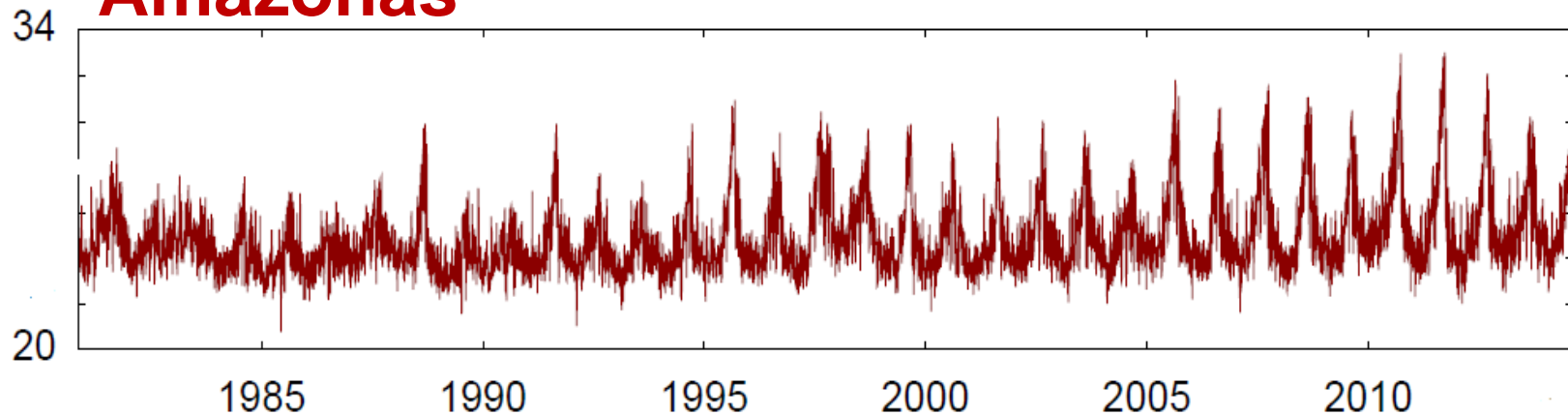
*G. Lancaster et al, “Surrogate data for hypothesis testing of physical systems”, Physics Reports 748, 1 (2018).*

# Relative decadal variations

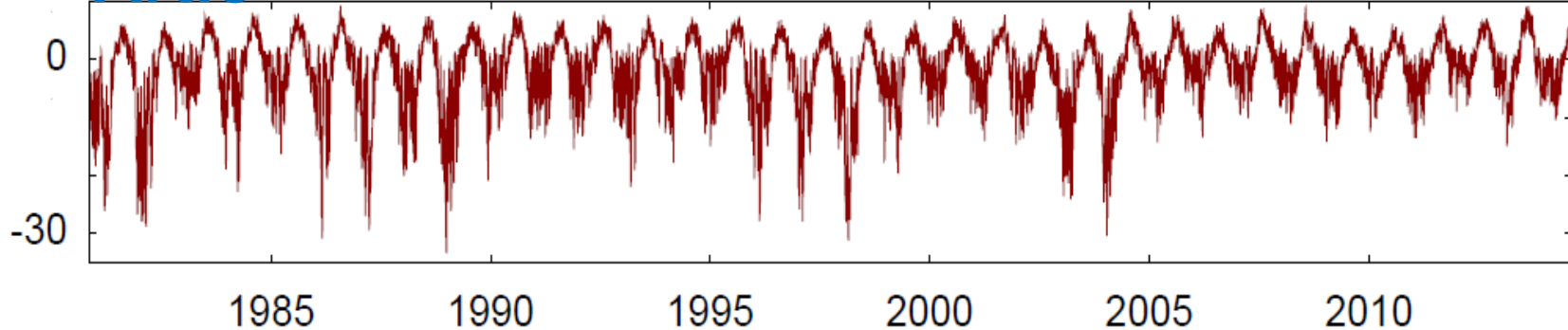


D. A. Zappala, M. Barreiro, C. Masoller, Earth Syst. Dynamics 9, 383 (2018)

# Amazonas



# Artic



- **Decrease of precipitation:** the solar radiation that is not used for evaporation is used to heat the ground.
- **Melting of sea ice:** during winter the air temperature is mitigated by the sea and tends to be more moderated.

# Outline

- Univariate analysis: Hilbert analysis

$$\{x_1, x_2, \dots, x_N\}$$

- Bivariate analysis: correlation and causality measures

$$\{x_1, x_2, \dots, x_N\} \quad \{y_1, y_2, \dots, y_N\}$$

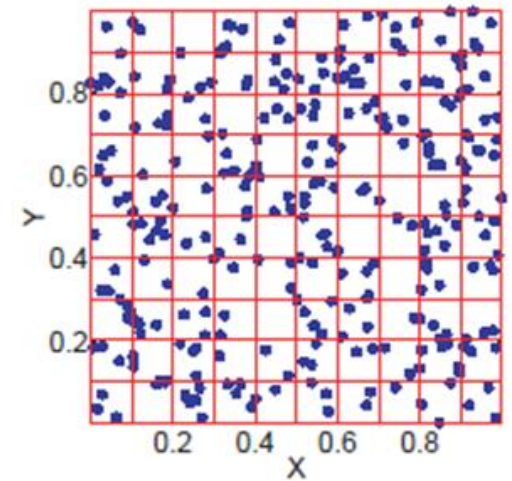
- Multivariate analysis: Ordinal analysis and complex networks

# Mutual Information (MI)

- *MI is calculated from probability distributions*

$$M_{ij} = \sum_{m,n} p_{ij}(m,n) \log \frac{p_{ij}(m,n)}{p_i(m)p_j(n)}$$

- It quantifies the reduction in uncertainty of one variable by knowing the other variable.
- *If  $X, Y$  are independent,  $MI = 0$ , else **MI > 0***
- For Gaussian processes:  $MI = -1/2 \log(1-\rho^2)$  where  $\rho$  is the cross-correlation coefficient.



# Transfer Entropy (TE) and Directionality Index (DI)

- TE: is the Conditional Mutual information, given the “past” of one of the variables.

$$TE(x,y) = MI(x, y|x_\tau)$$

$$TE(y,x) = MI(y, x|y_\tau)$$

- $MI(x,y) = MI(y,x)$  but  $TE(x,y) \neq TE(y,x)$
- Directionality Index:  $TE(x,y) - TE(y,x)$

*K. Hlaváčková-Schindler et al. / Physics Reports 441 (2007) 1–46*

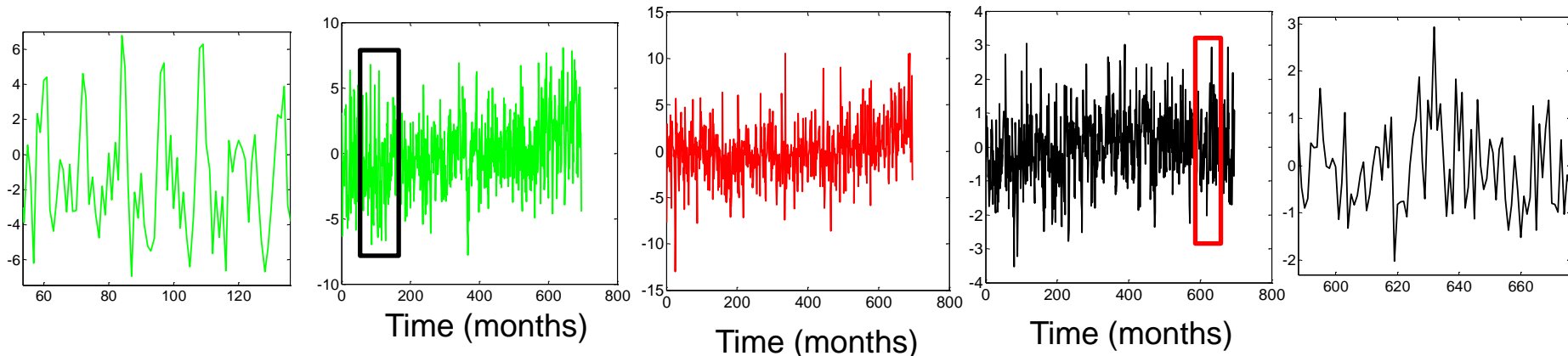
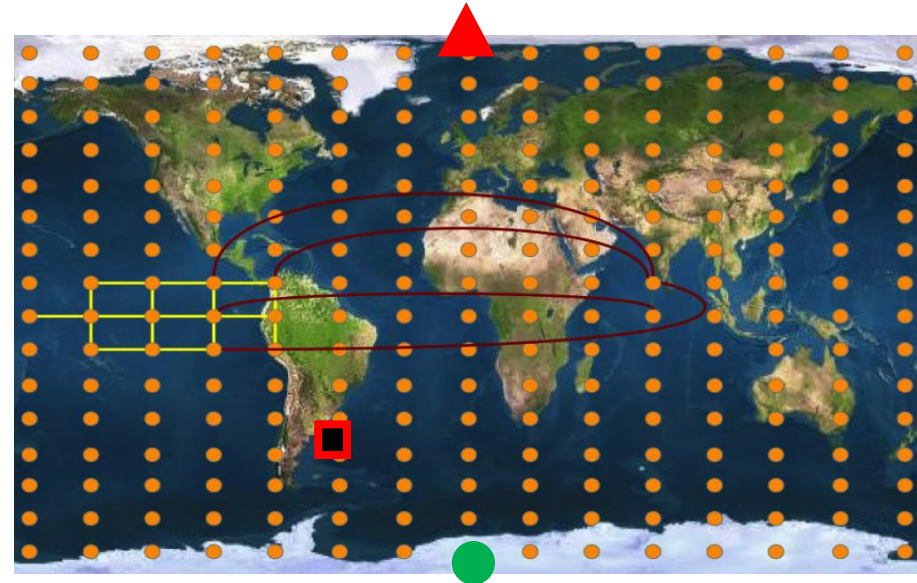
# Application: analysis of surface air temperature (SAT) anomalies.

**Anomaly** = annual solar cycle removed

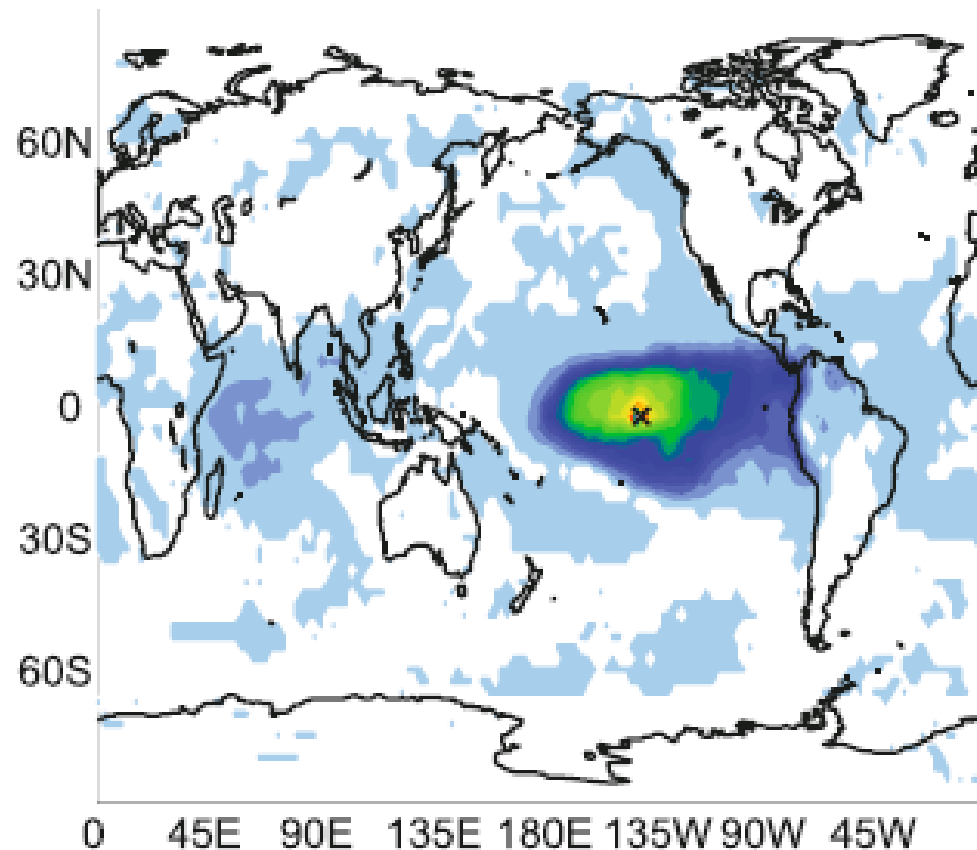
$2.5^\circ \times 2.5^\circ = 10226$  grid points

Monthly data  $\Rightarrow$  In each time series: 696 data points (1949-2006: 58 years x 12 months)

How does the data look like?

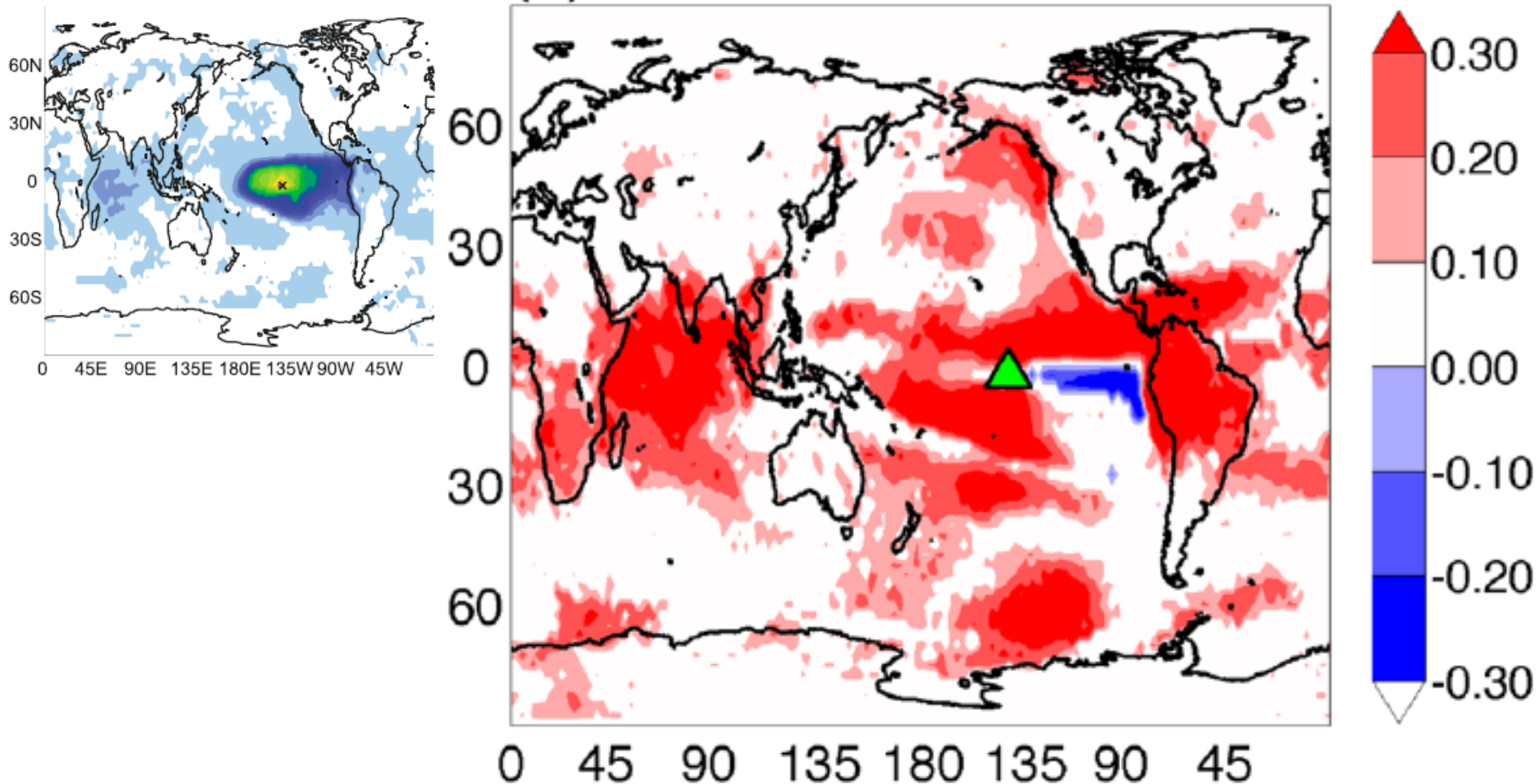


# Mutual Information of SAT anomaly in El Niño region and other regions (white: MI not significant)



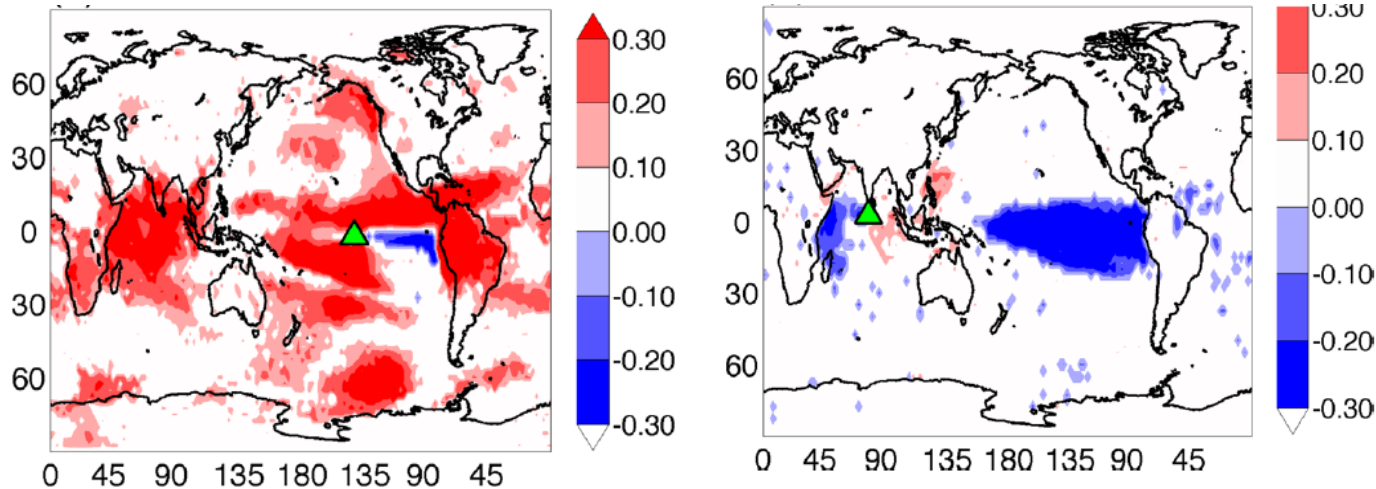
*J. I. Deza, M. Barreiro, C. Masoller, “Inferring interdependencies in climate networks constructed at inter-annual, intra-season and longer time scales”, Eur. Phys. J. ST 222, 511 (2013).*

# Directionality Index



J. I. Deza, M. Barreiro, and C. Masoller, “Assessing the direction of climate interactions by means of complex networks and information theoretic tools”, *Chaos* 25, 033105 (2015).

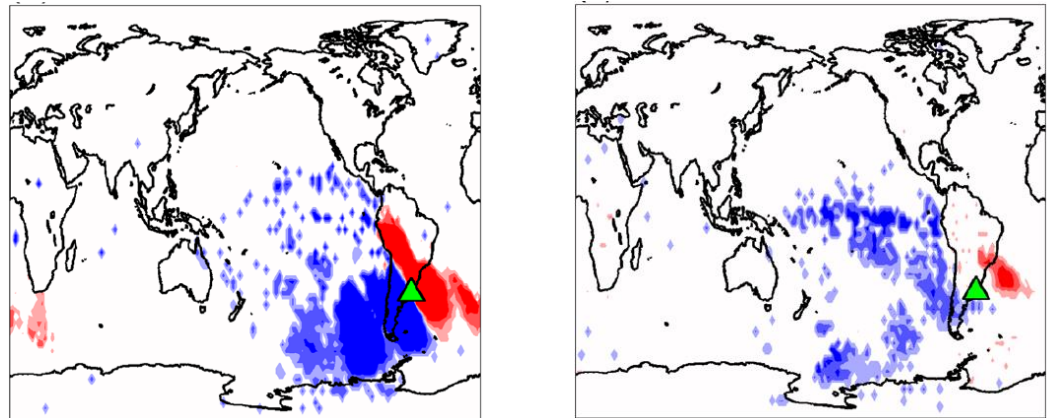
# Directionality Index



$$TE(x,y) = MI(x, y|x_\tau)$$

$$TE(y,x) = MI(y, x|y_\tau)$$

$$DI = TE(x,y) - TE(y,x)$$



J. I. Deza, M. Barreiro, and C. Masoller, Chaos 25, 033105 (2015).

# Problem: Transfer Entropy is computationally demanding

“simple” solution: use the expression that is valid for Gaussian distributions [  $MI = -1/2 \log(1-\rho^2)$  ]

Does this work? Check it out:

**scientific** reports

OPEN

**Fast and effective pseudo transfer entropy for bivariate data-driven causal inference**

Riccardo Silini<sup>✉</sup> & Cristina Masoller

<https://doi.org/10.1038/s41598-021-87818-3>



# Data Generating Processes and Performance Quantification

**Power: True Positives**

**Size: False Positives**

	Model
Y X	M <sub>0</sub>
	M <sub>1</sub>
	M <sub>2</sub>
Y → X	M <sub>3</sub>
	M <sub>4</sub>
	M <sub>5</sub>
	M <sub>6</sub>
	M <sub>7</sub>
	M <sub>8</sub>
	M <sub>9</sub>
	M <sub>10</sub>
	M <sub>11</sub>
	M <sub>12</sub>
Y ⇔ X	M <sub>13</sub>
	M <sub>14</sub>

$$x_t = (0.01 + 0.5 x_{t-1}^2)^{0.5} + E_{1t} \quad y_t = 0.5 y_{t-1} + E_{2t}$$

$$x_t = 0.6 x_{t-1} + 0.5 y_{t-1} + E_{1t} \quad y_t = 0.5 y_{t-1} + E_{2t}$$

$$x_t = 0.15 x_{t-1} + 0.7 y_{t-1} + E_{1t}$$

$$y_t = 0.1 y_{t-1} + 0.8 x_{t-1} + E_{2t}$$



# Results

		Model	pTE		
			$Y \rightarrow X$	$X \rightarrow Y$	
$Y \quad X$	{	M0	3.8	3.9	✓
		M1	2.3	2.6	
		M2	4.2	4.7	
$Y \rightarrow X$	{	M3	100	4.5	✓
		M4	80.7	3.8	
		M5	100	2.2	
		M6	100	1.8	
		M7	100	2.8	✗
		M8	100	4.5	
		M9	100	0.1	
		M10	62.6	3.1	
		M11	46.1	43.1	
$Y \Leftrightarrow X$	{	M12	99.9	1.0	✓
		M13	100	100	
		M14	100	100	



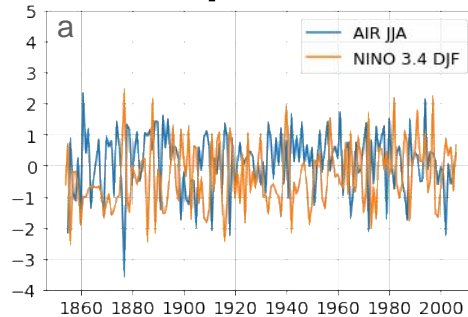
# Comparison with Granger Causality and Transfer Entropy

		Model	pTE		GC		TE		DI		
			$Y \rightarrow X$	$X \rightarrow Y$	$Y \rightarrow X$	$X \rightarrow Y$	$Y \rightarrow X$	$X \rightarrow Y$	pTE	GC	TE
$Y \rightleftharpoons X$	{	M0	3.8	3.9	5.1	5.0	4.4	4.4	-0.01	0.01	0.00
		M1	2.3	2.6	3.3	3.1	100	100	-0.06	0.03	0.00
		M2	4.2	4.7	5.5	5.9	4.7	4.9	-0.06	-0.04	-0.02
$Y \rightarrow X$	{	M3	100	4.5	100	4.8	70.2	5.6	0.91	0.91	0.85
		M4	80.7	3.8	84.2	4.9	96.0	4.7	0.91	0.89	0.91
		M5	100	2.2	100	3.1	100	3.8	0.96	0.94	0.93
		M6	100	1.8	100	2.8	100	4.3	0.96	0.95	0.92
		M7	100	2.8	100	3.4	100	4.0	0.95	0.93	0.92
		M8	100	4.5	100	5.6	100	100	0.91	0.89	0.00
		M9	100	0.1	100	0.1	100	100	1.00	1.00	0.00
		M10	62.6	3.1	67.3	4.3	12.2	4.5	0.91	0.88	0.46
		M11	46.1	43.1	53.1	49.8	37.8	45.0	0.03	0.03	-0.09
		M12	99.9	1.0	100	0.9	100	0	1.0	1.0	1.0
$Y \rightleftharpoons X$	{	M13	100	100	100	100	100	100	0.00	0.00	0.00
		M14	100	100	100	100	100	100	0.00	0.00	0.00

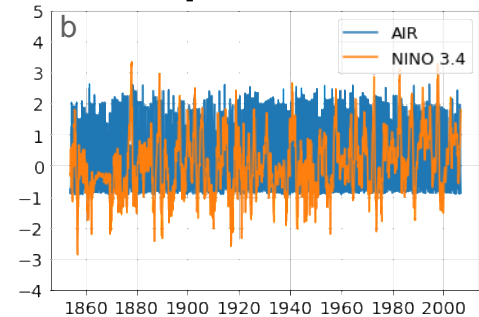


# Application to real data $\text{NINO3.4} \leftrightarrow \text{All India Rainfall}$

Yearly  
sampled (152)



Monthly  
sampled (1836)



IAAF

pTE

GC

TE

$\text{NINO3.4} \rightarrow \text{AIR}$

0.04 s

$\text{NINO3.4} \rightarrow \text{AIR}$

0.4 s

$\text{NINO3.4} \leftrightarrow \text{AIR}$

1 s

$\text{NINO3.4} \leftarrow \text{AIR}$

0.5 s

$\text{NINO3.4} \leftarrow \text{AIR}$

0.9 s

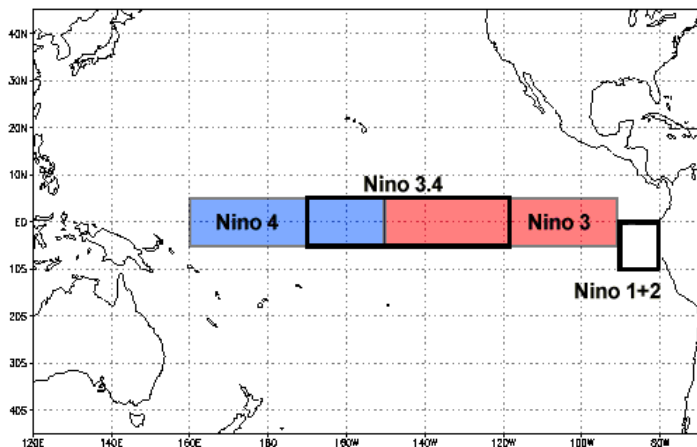
$\text{NINO3.4} \leftrightarrow \text{AIR}$

40<sub>9</sub>  
3  
68 s

# How much time can we save?

For two time-series of 500 data points (1 data point per month, 40 years):

TE: **112 ms** but pTE: **4 ms**



8000 grid points (high resolution)  
⇒  $64 \times 10^6$  pairs

⇒ **829 days** (TE) vs. **29 days** (pTE).

(without “surrogate” analysis)

But, there is a price to pay, no “free lunch”.

<https://github.com/riccardosilini/pTE>

# Outline

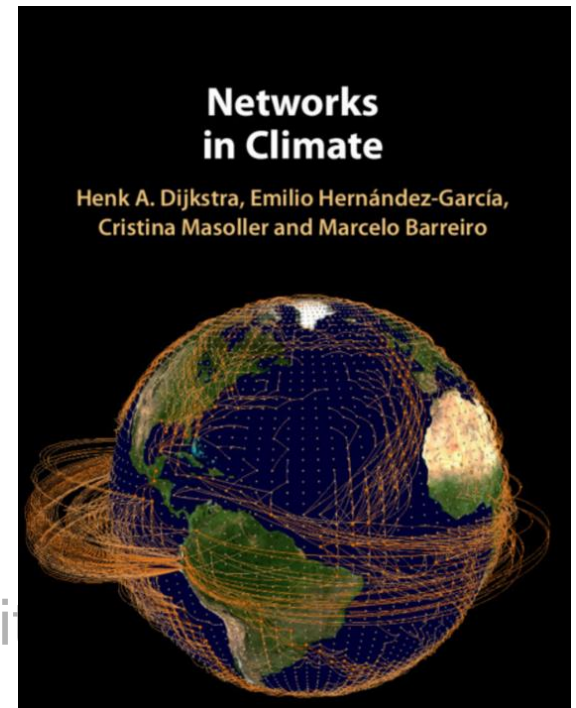
- Univariate analysis: Hilbert analysis

$$\{x_1, x_2, \dots, x_N\}$$

- Bivariate analysis: correlation and causality

$$\{x_1, x_2, \dots, x_N\} \quad \{y_1, y_2, \dots, y_N\}$$

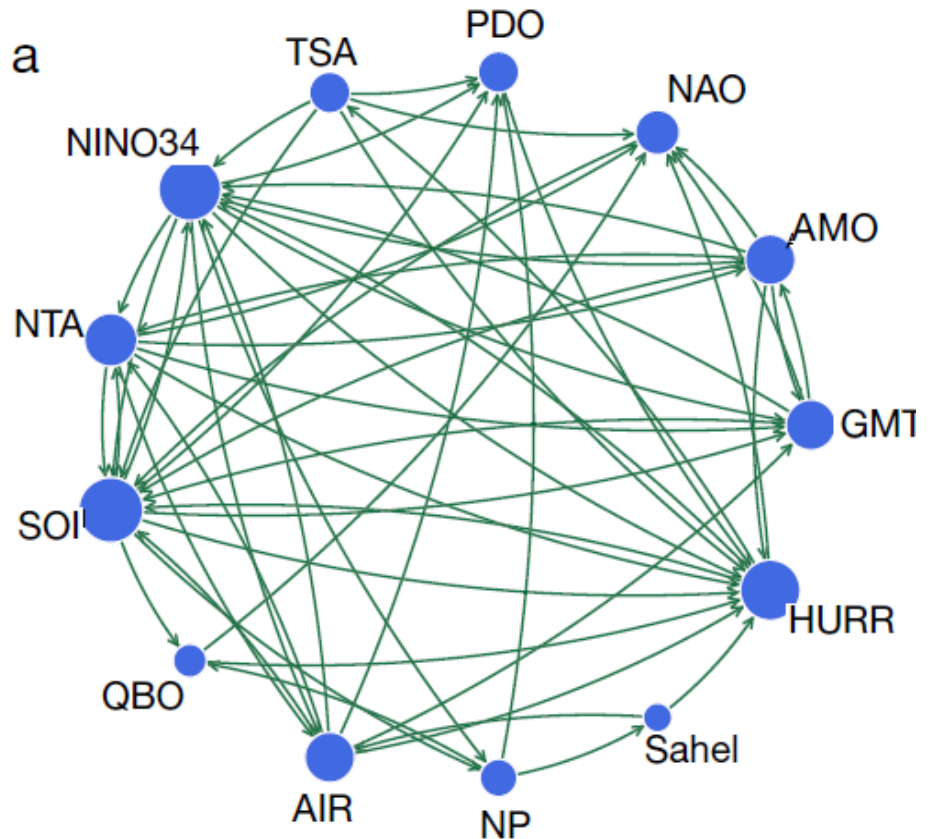
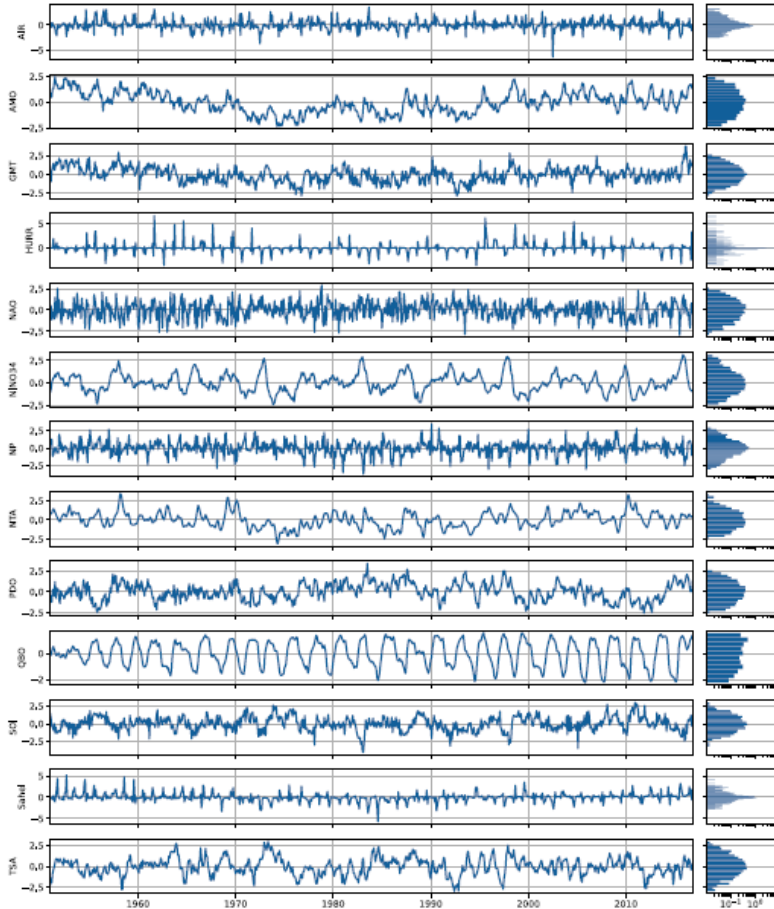
- Multivariate analysis: Ordinal analysis and complex networks



Cambridge University Press 2019

# Directed network of climatic indices

Constructed calculated pTE with different lags



R. Silini, G. Tirabassi, M Barreiro, L. Ferranti, C. Masoller, “Assessing causal dependencies in climatic indices”, *Climate Dynamics* 61, 79–89 (2023).

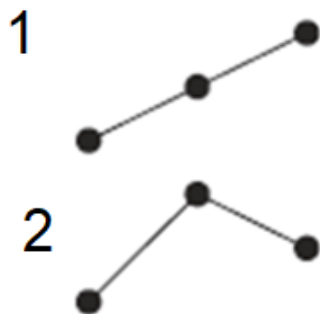
**Ordinal analysis:**  
**A nonlinear way to select the**  
**time-scale of the analysis**

# Ordinal analysis

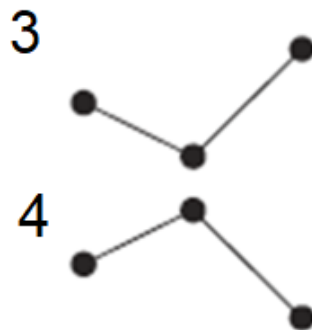
$$\{\dots x_i, x_{i+1}, x_{i+2}, \dots\}$$

How can three data points (let's say 2, 5, 7) be ordered?

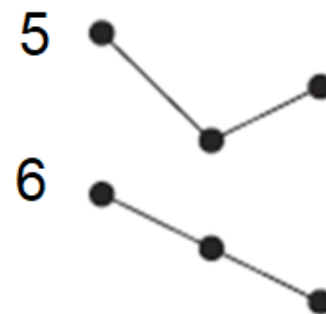
$\{\dots 2, 5, 7 \dots\}$



$\{\dots 5, 2, 7 \dots\}$



$\{\dots 7, 2, 5 \dots\}$



$\{\dots 2, 7, 5 \dots\}$

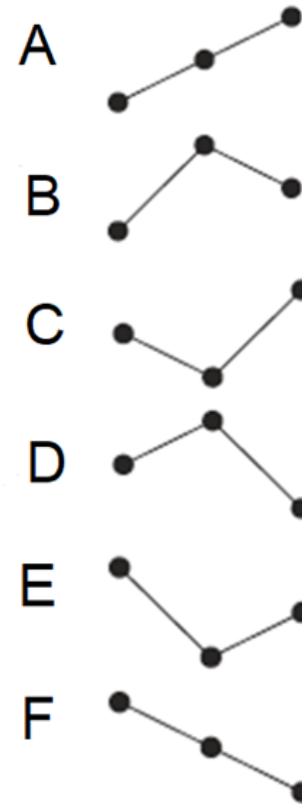
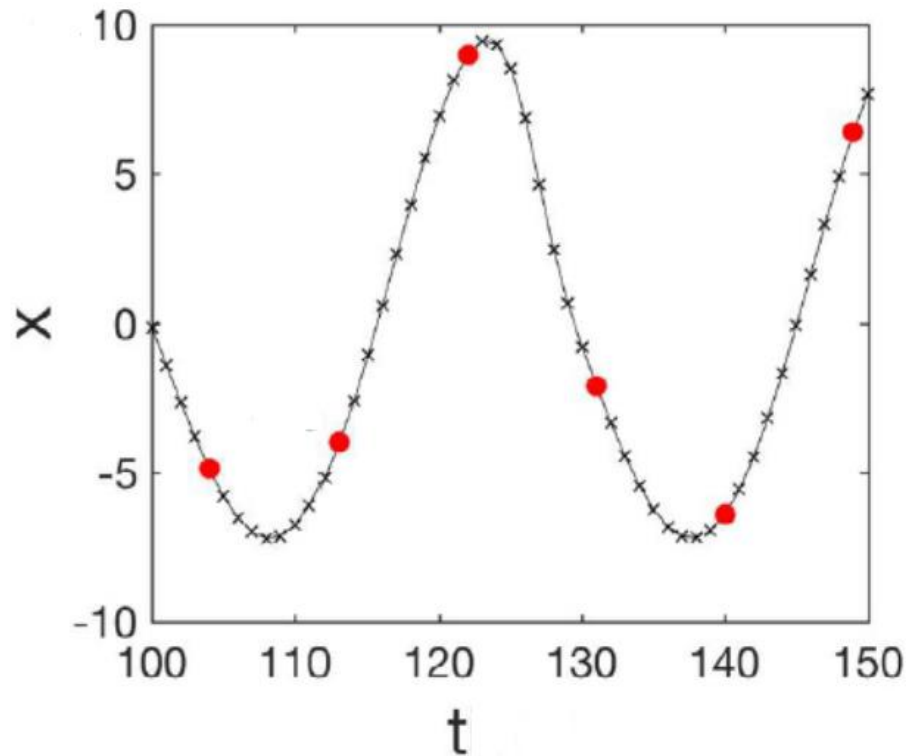
$\{\dots 5, 7, 2 \dots\}$

$\{\dots 7, 5, 2 \dots\}$

Bandt and Pompe: Phys. Rev. Lett. 2002

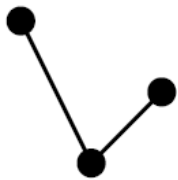
Which is the “message” “encoded” in the red dots?

**{A, B, F, C}**

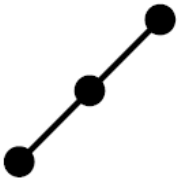


Using **lagged points** to define the patterns allows to select the time scale of the analysis, useful for seasonal data

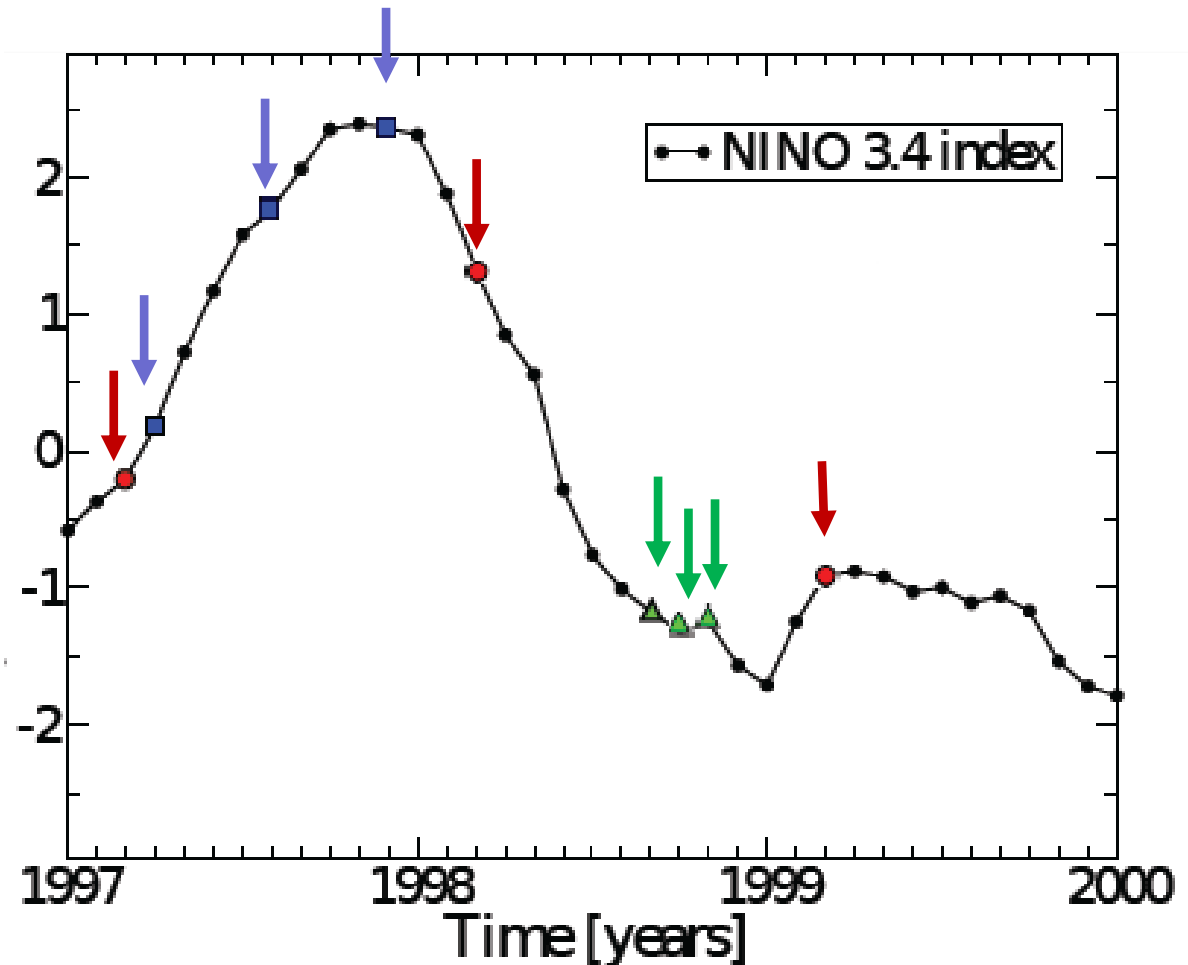
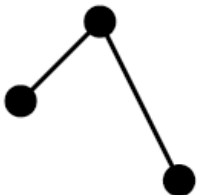
Intra-season



Intra-annual



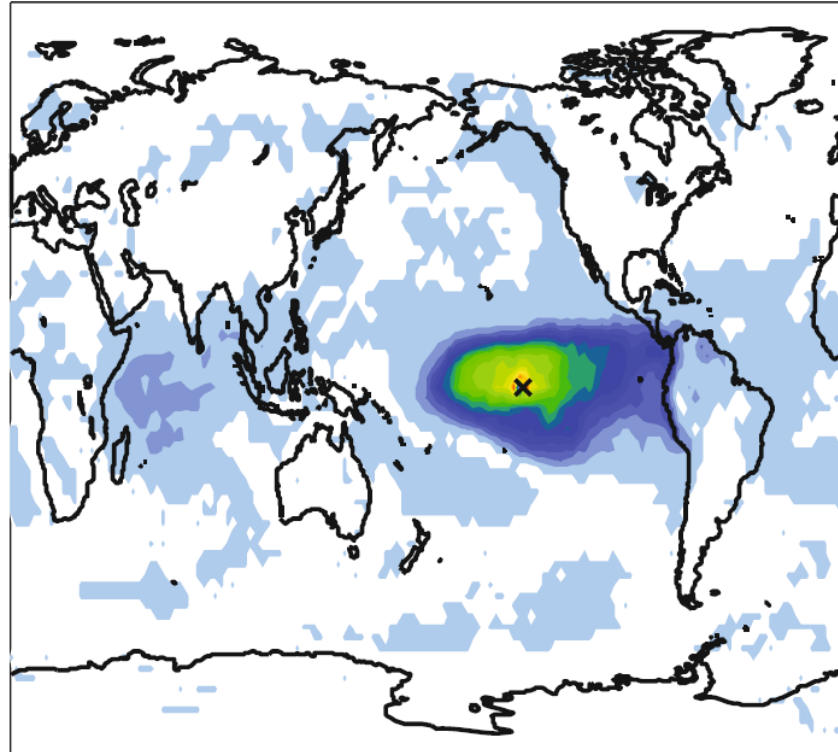
Inter-annual



# Mutual Information of SAT anomaly in El Niño region and other regions (**shown before**, white: MI not significant)

$$M_{ij} = \sum_{m,n} p_{ij}(m,n) \log \frac{p_{ij}(m,n)}{p_i(m)p_j(n)}$$

MI from  
probabilities  
of SAT  
values

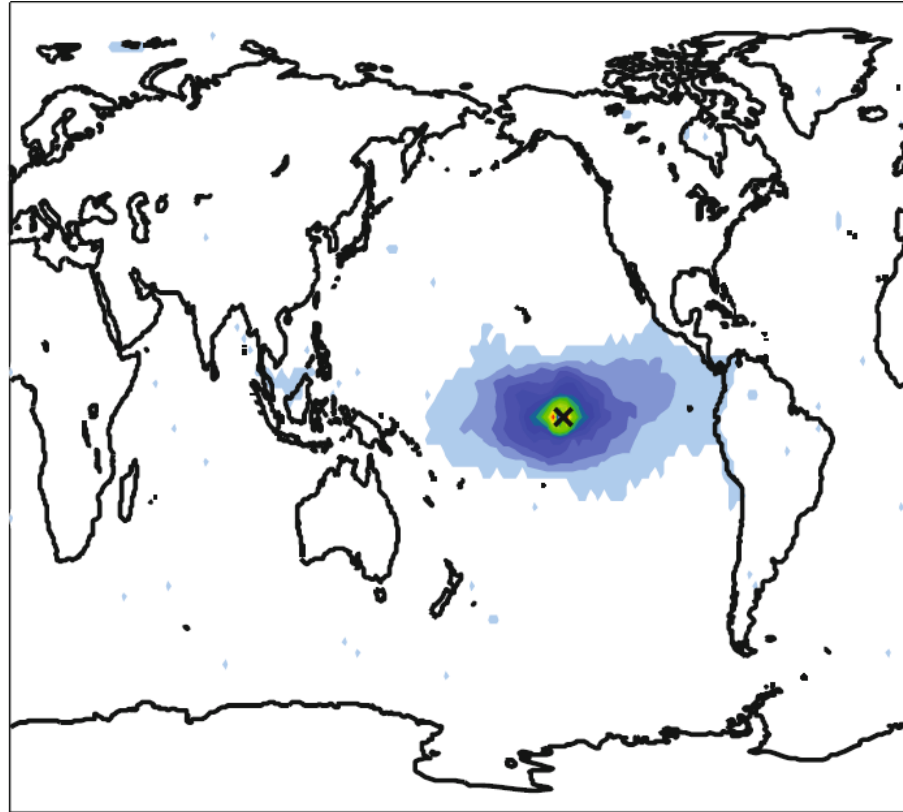


*J. I. Deza, M. Barreiro, C. Masoller, Eur. Phys. J. ST 222, 511 (2013)*

# Mutual Information (color code) from probabilities of ordinal patterns (white: MI not significant)

$$M_{ij} = \sum_{m,n} p_{ij}(m,n) \log \frac{p_{ij}(m,n)}{p_i(m)p_j(n)}$$

MI from probabilities of ordinal patterns defined by values in 3 consecutive months.

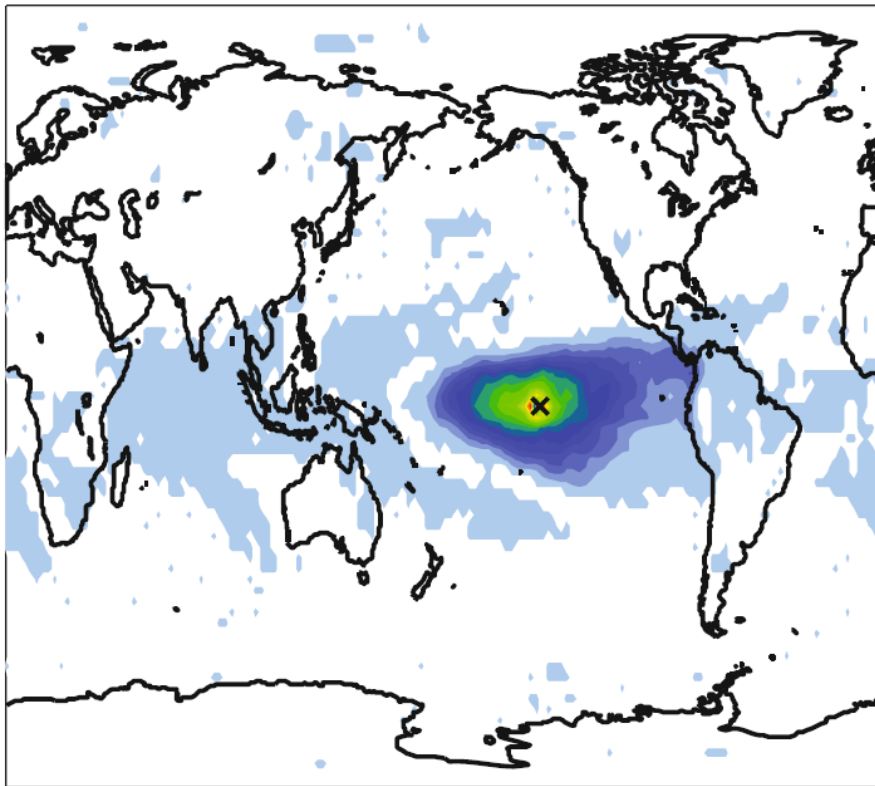


*J. I. Deza, M. Barreiro, C. Masoller, Eur. Phys. J. ST 222, 511 (2013).*

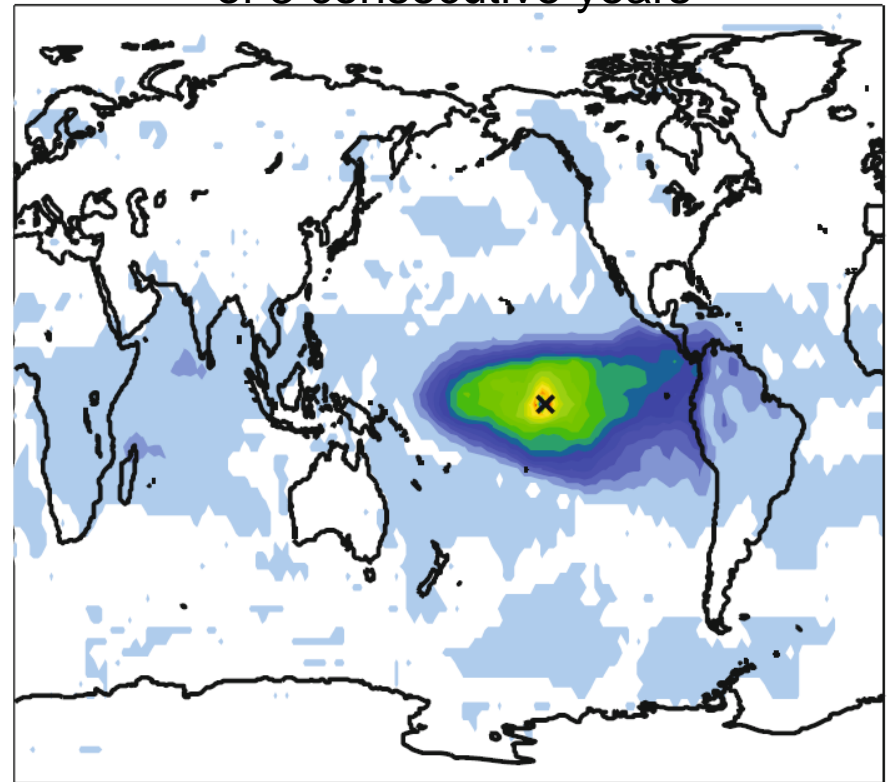
# Comparison of two ways to calculate the ordinal probabilities, used to calculate the mutual information

$$M_{ij} = \sum_{m,n} p_{ij}(m,n) \log \frac{p_{ij}(m,n)}{p_i(m)p_j(n)}$$

Patterns defined by 3 values in a year



Patterns defined by data values of 3 consecutive years

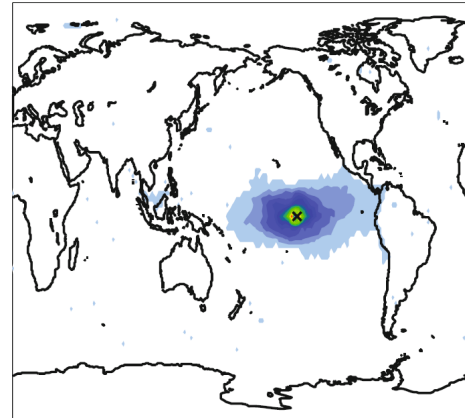
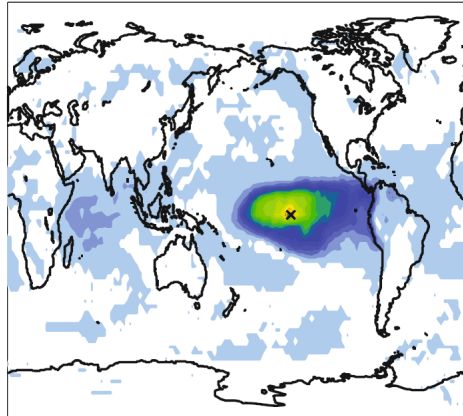


*J. I. Deza, M. Barreiro, C. Masoller, Eur. Phys. J. ST 222, 511 (2013).*

# Comparison

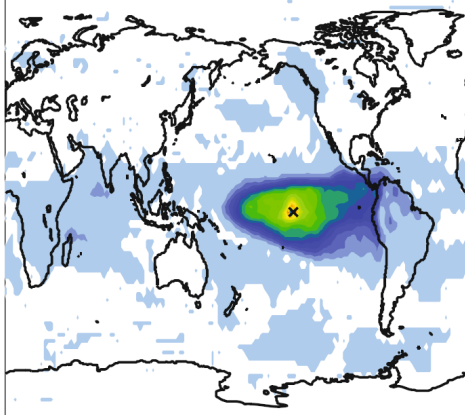
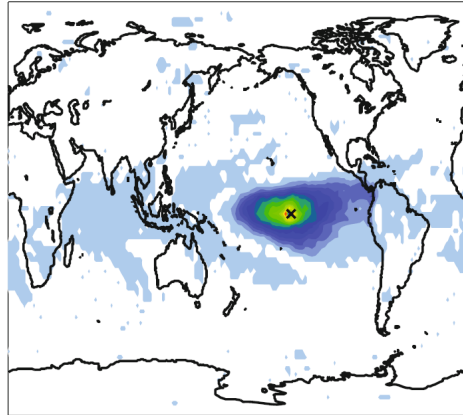
$$M_{ij} = \sum_{m,n} p_{ij}(m,n) \log \frac{p_{ij}(m,n)}{p_i(m)p_j(n)}$$

probabilities of  
SAT values



probabilities of  
ordinal patterns  
defined by values  
in 3 consecutive  
months.

probabilities of  
patterns defined  
by 3 values in a  
year.



probabilities of  
patterns defined  
by values in 3  
consecutive  
years.

*J. I. Deza, M. Barreiro, C. Masoller, Eur. Phys. J. ST 222, 511 (2013).*

# Take home messages

- Hilbert analysis and ordinal analysis are versatile tools that can provide new insights into climate phenomena.
- Mutual information and other information-based measures can be calculated in terms of the probabilities of ordinal patterns, allowing to select the time-scale of the analysis.
- Different large-scale spatial structures are uncovered when using different lags between the data points that define the ordinal patterns.

# Thank you for your attention

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